



# The Radiative Feedback of the Tropical Anvil Clouds: Negative or Positive?

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CERES Science Team Meeting  
Newport News, VA, May 2, 2001

# Acknowledgement

- David Young, Tak Wong, and David Kratz gave many suggestions and comments on this study.
- Ed Kizer helped to set up ERBE-like codes.
- Erika Geier, Seiji Kato, Norman Loeb, and Alice Fan provided the support in SSF data processing.



# Outline

## 1. Background

Lindzen et al.' (2001) climate feedback:  
observation & model

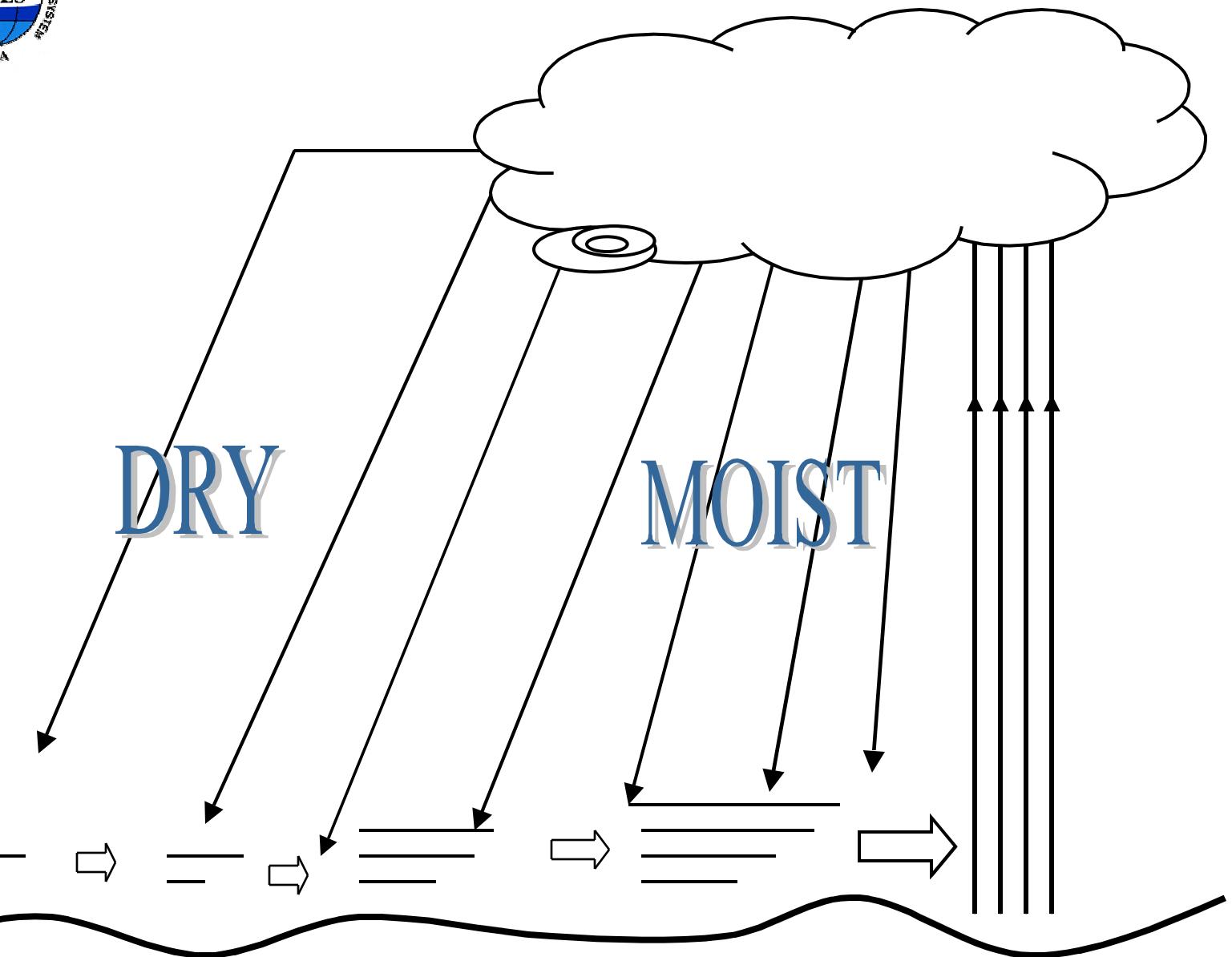
## 2. CERES Data

## 3. 3.5-box Model Calculation

## 4. Perturbation Analysis

## 5. Summary

# 1. Background

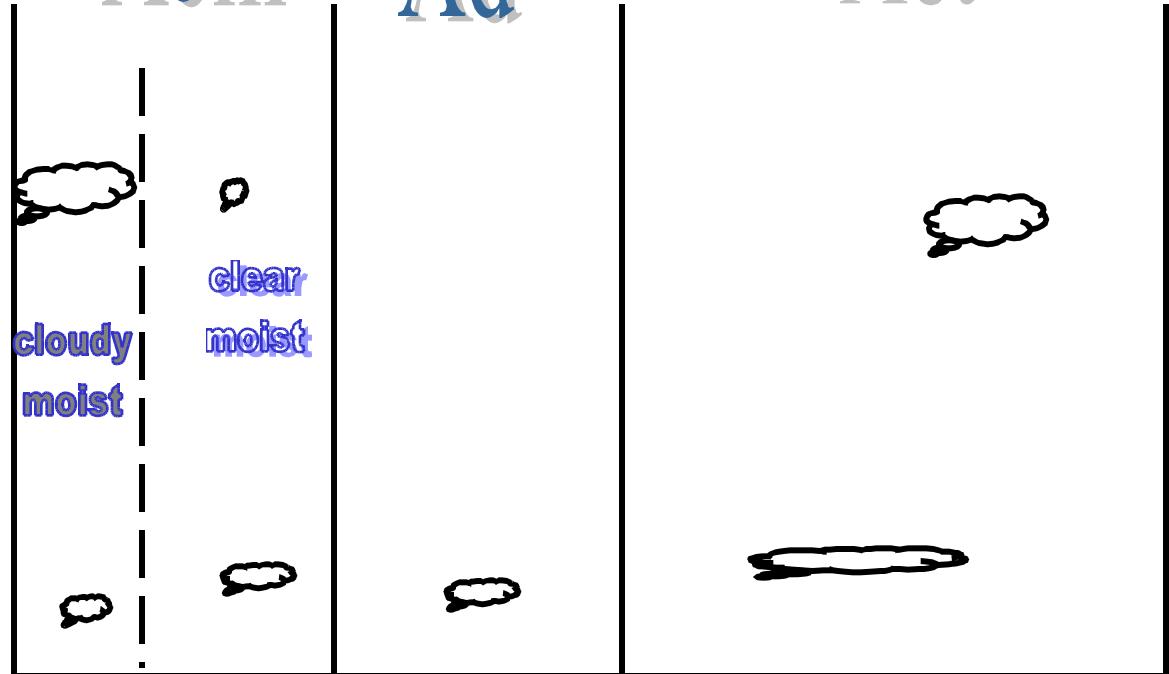


**Atmospheric Moisturization**

## Tropics

Acm

Ad



$$T_{st} = T_s + 10K$$

$$T_{set} = T_s - 10K$$



## 1. Background (cont.)

Based on the anvil variations with SST observed from GMS data and 3.5-box greenhouse model, Lindzen et al. (2001) proposed a very strong negative radiative feedback of the clouds on climate change ( $-0.45 \sim -1.1$ ; or **IR iris**).

Note: cloud amount with SST      c.f.  
Chambers et al.



# 1. Background (cont.): main points

Q: Do CERES data show the similar cloud change with SST, and feedback processes?

(Since we do not know where many values in Lindzen et al. come from)

1. Confirms the anvil clouds over oceans decrease with increase SST with the rate.
2. Large differences in CERES observed SW & LW fluxes from those of Lindzen et al.
3. Significant differences in calculated cloud feedback from those Lindzen et al.' results.



## 2. CERES Data

CERES/TRMM ERBE-like observations:  
the area coverage of tropical dry, clear moist,  
and cloudy moist regions, albedo, incoming  
shortwave (SW ) and outgoing longwave  
(LW ).

SSF Data: Tb( $10.8\mu\text{m}$ ) + broadband LW &  
SW measurements ( $\pm 30^\circ\text{N}$ ; 01 ~ 08, 1998)



## 2. CERES Data (cont.)

Definitions of clouds & climate regimes:

convective clouds:  $Tb(10.8) < 220K$

cloudy moist:  $Tb(10.8) < 260K$  (anvil)

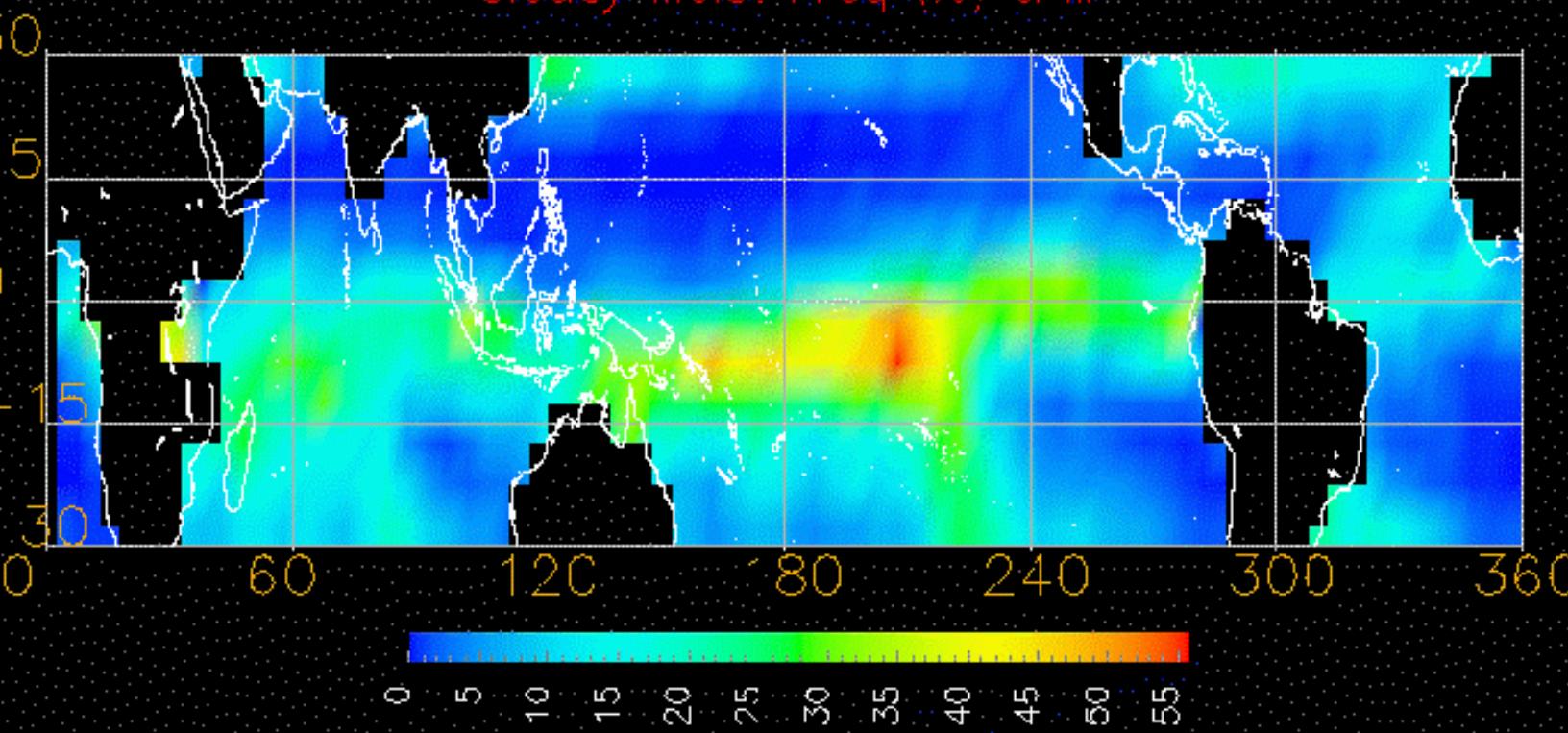
dry area: broadband LW  $> LW50$

LW50: 50% percentile of 8-month LW statistics

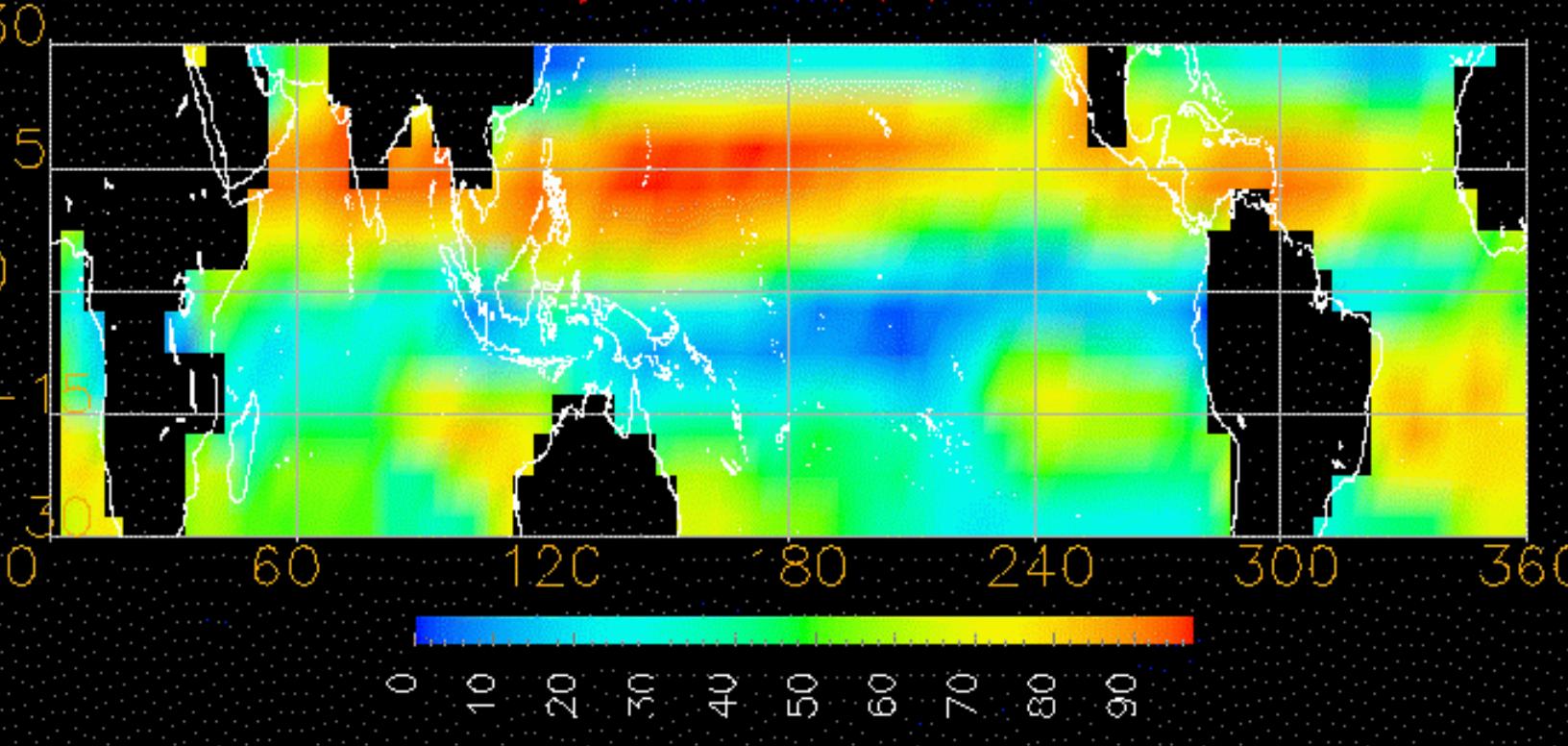
clear moist: all other pixels



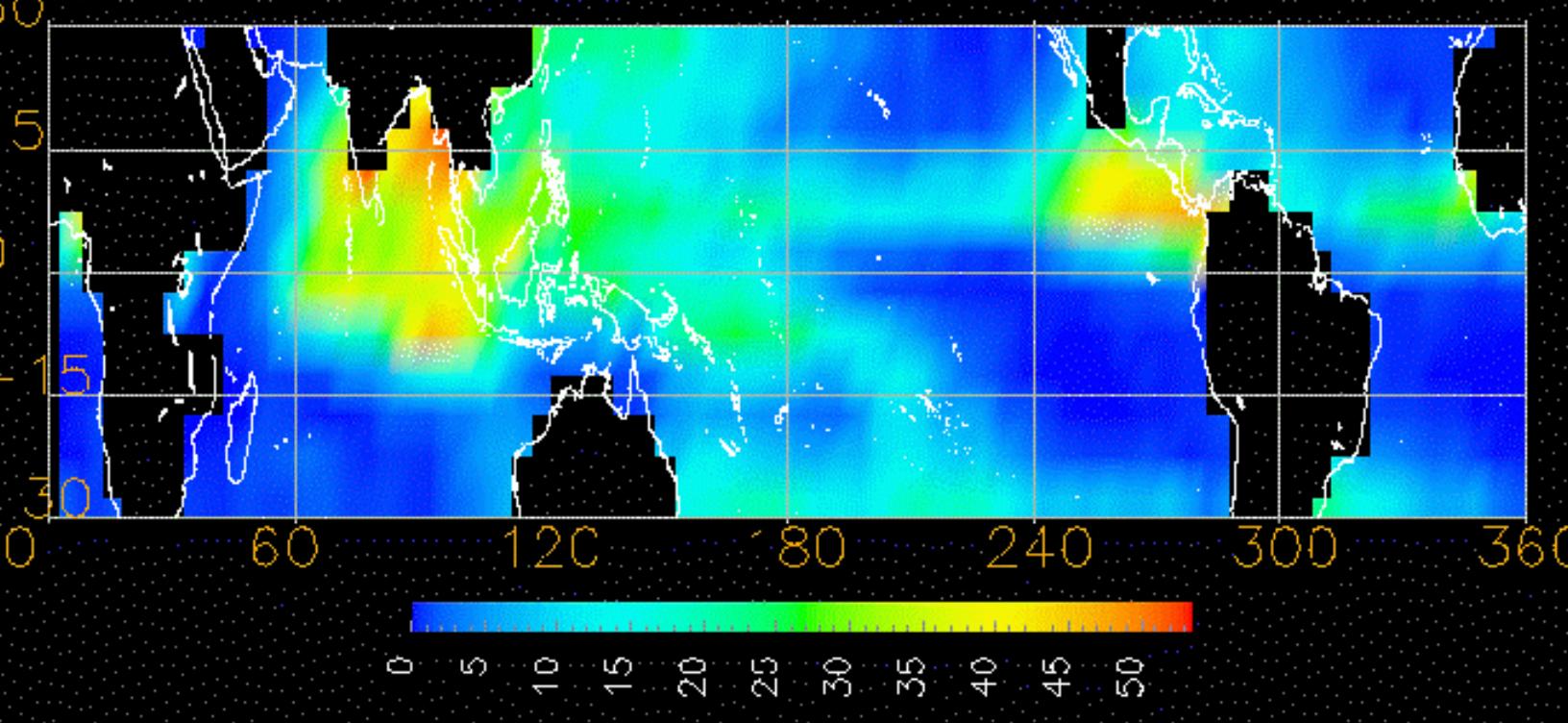
Cloudy Mois Freq (%) JFM



Dry Area Freq (%) JFM

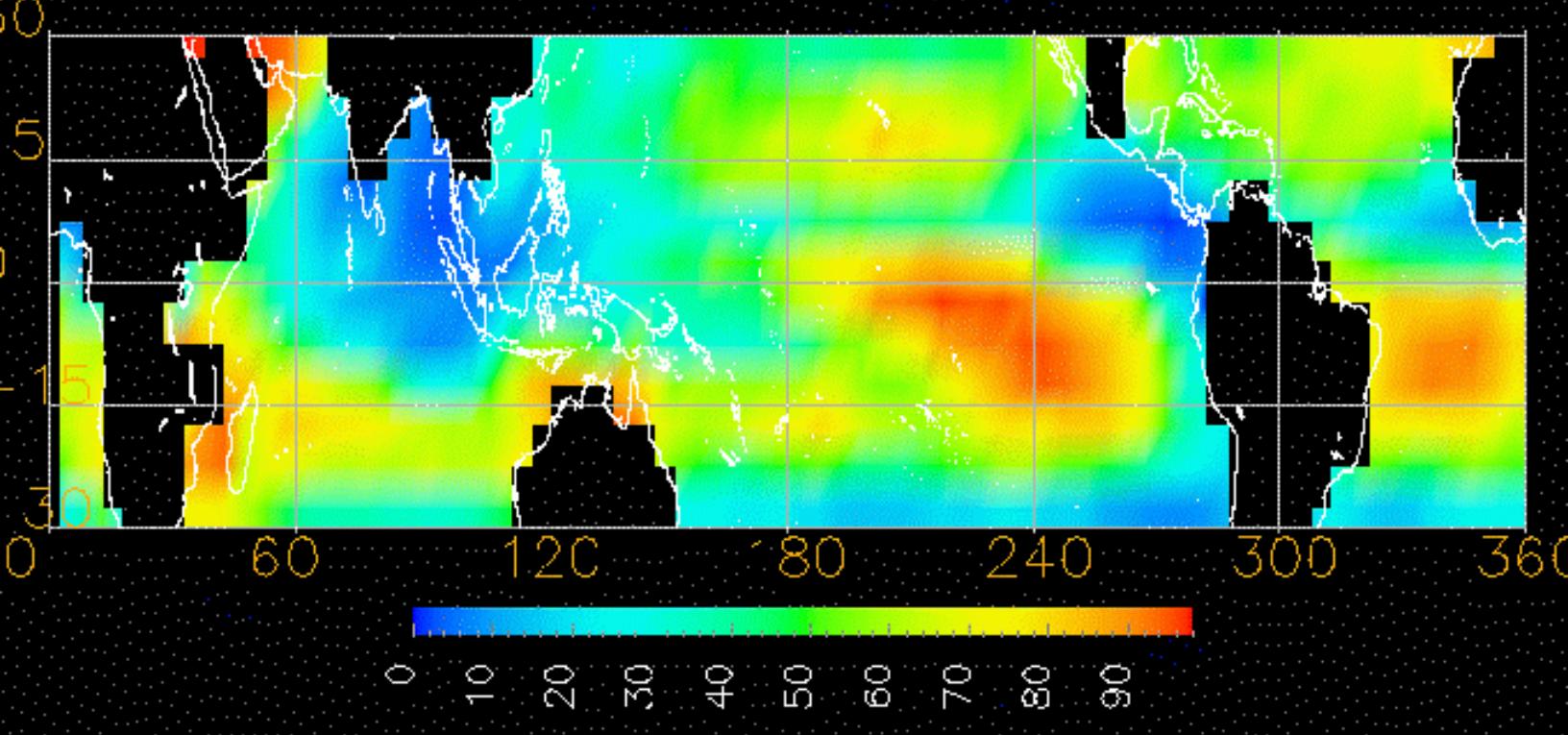


Cloudy Mois. Freq. (%) JJA





Dry Area Freq (%) JJA



## 2. CERES Data (cont.)



	LaRC	CERES		Lindzen et al.		
	dry	clear moist	cloudy moist	dry	clear moist	cloudy moist
q	0.5	0.4	0.1	0.5	0.28	0.22
edo	0.154	0.258	0.510	0.211	0.211	0.349
	338.7	297.1	196.2	315.9	315.9	260.6
	287.7	253.9	154.8	303.1	263.1	137.7



### 3. 3.5-box Model Calculation

$$A_{clrm}SW_{clrm} + A_{cldm}SW_{cldm} + A_dSW_d + A_{et}SW_{et} = Q_0(1 - \sigma_0) = (A_{clrm}T_{eclrm}^4 + A_{cldm}T_{ecldm}^4 + A_dT_{ed}^4 + A_{et}T_{eet}^4) \quad (1)$$

Note: LW = T<sub>e</sub><sup>4</sup> = (T<sub>s</sub> - T<sub>e</sub>)<sup>4</sup>

clr clear; cld cloudy; m moist; d dry;  
t tropical; et extratropical; e emission;



## 3.5-box greenhouse model (cont.)

$$A_{cldm} = A_{cm}^\circ (1 + \mu);$$

$$A_{cm} = A_{cm}^\circ (1 + \mu);$$

$$A_{clrm} = A_{cm} - A_{cldm}; \quad A_d = A_t - A_{cm};$$

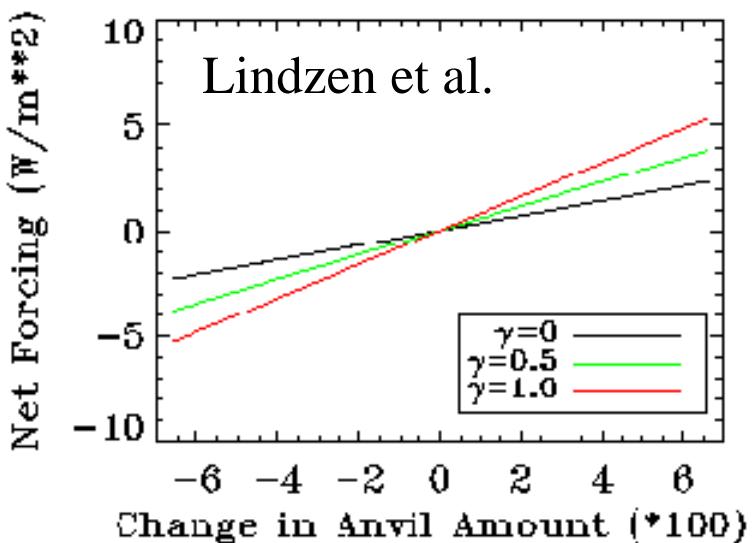
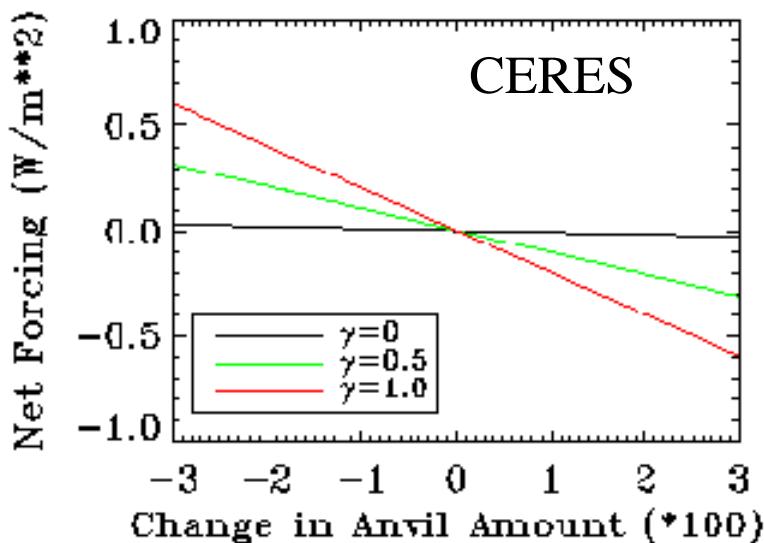
= 0 corresponding cloudy moist to clear moist

changes

> 0    3-way, i.e., clear moist, cloudy moist, and dry regions, changes, which produces stronger feedback.



# radiative forcing: IR iris?



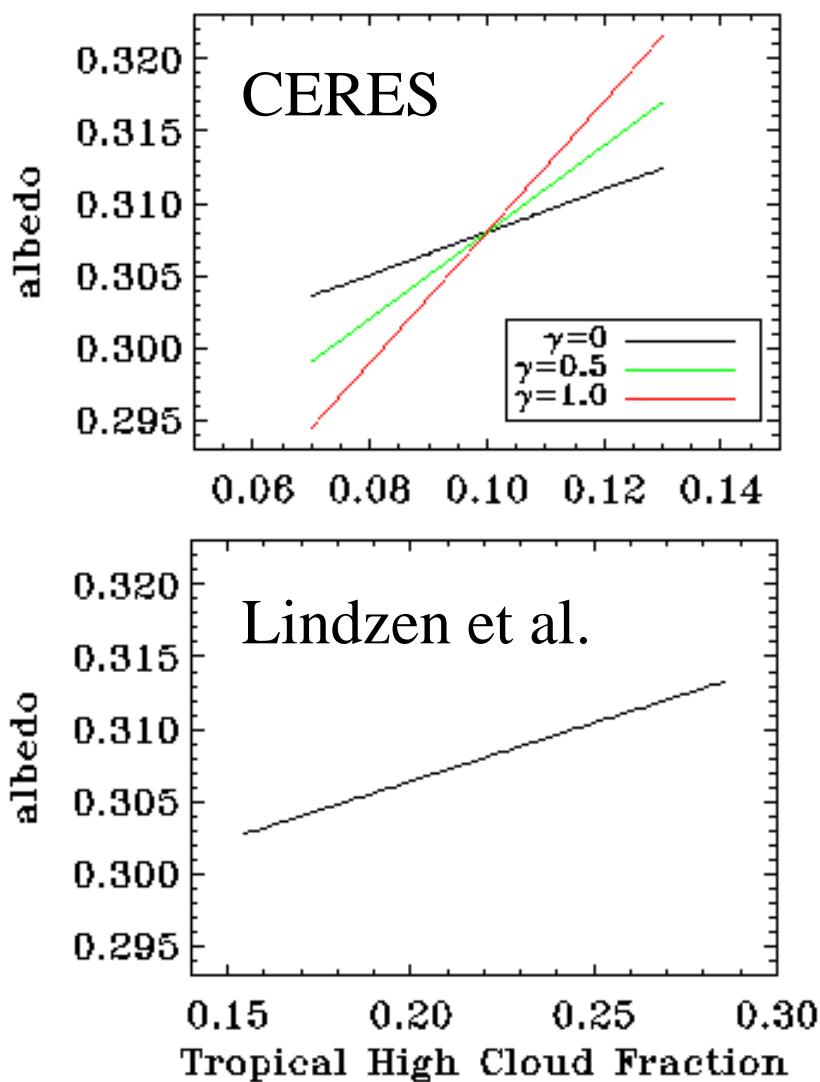
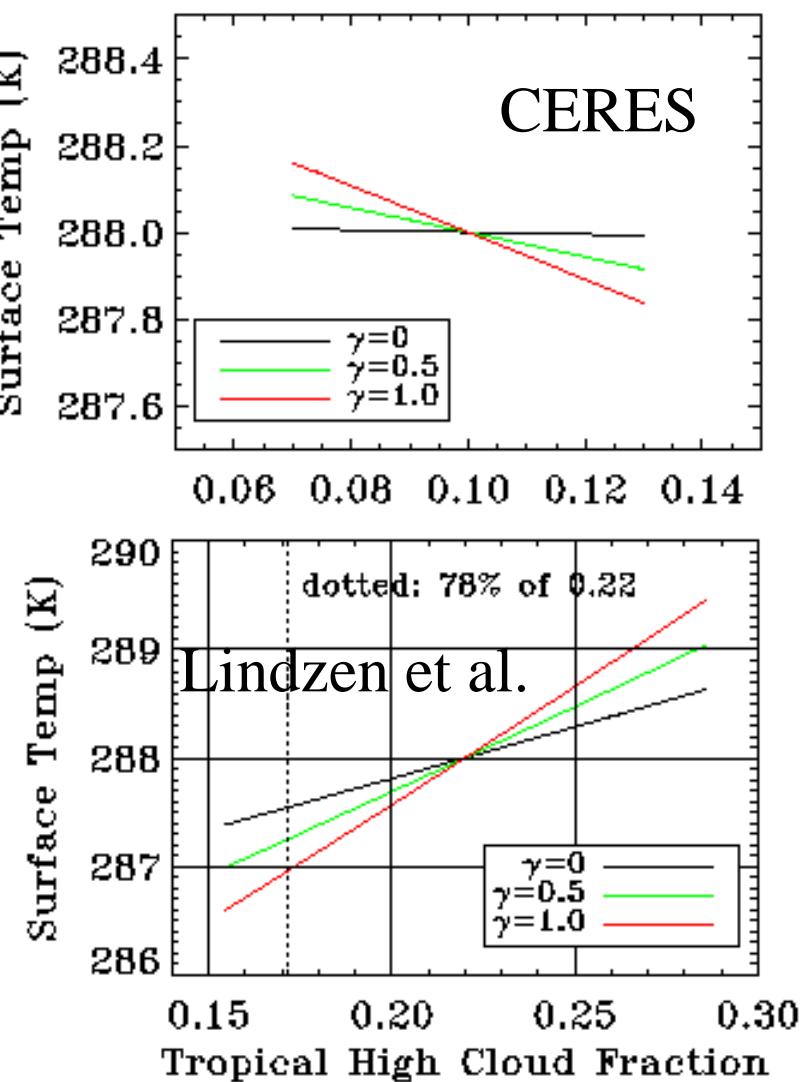
$$\begin{aligned} \text{Global Net Forcing} &= \text{SW} - \text{LW} \\ C = CC^\circ + CC \text{ and } Ts = T^\circ s &\quad CC = \text{change in anvil amount} \end{aligned}$$

anvil clouds  $\downarrow$  global albedo  $\downarrow$  & SW  $\uparrow$

Lindzen et al.: LW  $\nearrow$  the net radiation  $\downarrow$

CERES: NOT seen!

# Simulated cloud feedback





## 4. Perturbation Analysis

$$A_{clrm} SW_{clrm} + A_{cldm} SW_{cldm} + A_d SW_d + A_{et} SW_{et} = Q_0(1 - \alpha_0) = \\ (A_{clrm} T^4_{eclrm} + A_{cldm} T^4_{ecldm} + A_d T^4_{ed} + A_{et} T^4_{eet}) \dots \dots \quad (1)$$

$$A_{clrm} SW_{clrm} + A_{cldm} SW_{cldm} + A_d SW_d = -Q_0 \alpha_0 = \\ (A_{clrm} T^4_{eclrm} + A_{cldm} T^4_{ecldm} + A_d T^4_{ed}) + \\ 4 (A_{clrm} T^3_{eclrm} + A_{cldm} T^3_{ecldm} + A_d T^3_{ed} + A_{et} T^3_{eet}) \quad Ts \dots \quad (2)$$

$$A_{clrm} + A_{cldm} + A_d = A_t; \quad A_{clrm} + A_{cldm} + A_d = 0 \dots \dots \quad (3)$$

$$4 (A_{clrm} T^3_{eclrm} + A_{cldm} T^3_{ecldm} + A_d T^3_{ed} + A_{et} T^3_{eet}) \quad Ts = \\ A_{clrm} ((SW_{clrm} - LW_{clrm}) - (SW_d - LW_d)) + \\ A_{cldm} ((SW_{cldm} - LW_{cldm}) - (SW_d - LW_d)) \dots \dots \quad (4)$$



## 4. Perturbation Analysis (cont.)

	LaRC — CERES			Lindzen et al.		
	Dry	Clear Moist	Cloudy Moist	Dry	Clear Moist	Cloudy Moist
o	<b>0.5</b>	<b>0.4</b>	<b>0.1</b>	<b>0.5</b>	<b>0.28</b>	<b>0.22</b>
o	<b>0.154</b>	<b>0.258</b>	<b>0.510</b>	<b>0.211</b>	<b>0.211</b>	<b>0.349</b>
	<b>338.7</b>	<b>297.1</b>	<b>196.2</b>	<b>315.9</b>	<b>315.9</b>	<b>260.6</b>
	<b>287.7</b>	<b>253.9</b>	<b>154.8</b>	<b>303.1</b>	<b>263.1</b>	<b>137.7</b>
Radiation	<b>51.0</b>	<b>43.2</b>	<b>41.4</b>	<b>12.8</b>	<b>52.8</b>	<b>122.9</b>
net versus net		<b>-7.8</b>	<b>-9.6</b>		<b>40.0</b>	<b>110.1</b>
dry moist vs. moist			<b>-1.8</b>			<b>70.1</b>



## 4. Perturbation Analysis (cont.)

$$\text{CERES: } 3.707 \quad T_s = -7.8 \quad A_{clrm} - 9.6 \quad A_{cldm} \quad \dots \quad (5-1)$$

$$\text{Lindzen et al. } 3.703 \quad T_s = 40.0 \quad A_{clrm} + 110 \quad A_{cldm} \quad \dots \quad (5-2)$$

1).  $A_d = \text{Const}$  ( $= 0$ ) no dry region terms in Eq. 4.

$$\text{And } A_{clrm} = -A_{cldm}$$

CERES only  $-1.8 \quad A_{cldm}$ ,  
 $T_s = 0.01K$

Lindzen et al.  $70.1 \quad A_{cldm}$   
 $T_s = -0.46K$

2)  $A_{cm}$  increasing rate =  $A_{cldm}$  increasing rate ( $= 1$ )

the same  $A_{clrm}$  increasing rate much stronger feedback

$T_s = 0.12K$   $T_s = -1.05K$

These sensitivities or feedback factors are the same as those of previous model simulations (c.f. Figures)



## SW & LW consistency

Fu-Liou ice clouds: for Lindzen et al.' values

$H_c = 15\text{ km}$  (upper limit for warming)

With  $LW = 138 \text{ W/m}^2$        $r_{cldm} = 0.39$

if add  $=15 \sim 50$  into distribution,  $r_{cldm}$  doubled  
with small change in  $LW$  .

Increased SW cooling effects: e.g,

$=16$ ,  $r_{cldm} = 0.63$ ,     $LW = -35 \text{ W/m}^2$  (warming)

$SW = (0.63 - 0.4)400 = 92 \text{ W/m}^2$  (cooling)

# SW & LW consistency (cont.)

Fu-Liou model      Cloudy Moist

$\tau$	freq	albedo	LW $\downarrow$
1	<b>0.15</b>	<b>0.261</b>	<b>207.4</b>
2	<b>0.15</b>	<b>0.356</b>	<b>173.8</b>
4	<b>0.15</b>	<b>0.448</b>	<b>148.8</b>
8	<b>0.15</b>	<b>0.549</b>	<b>139.8</b>
20	<b>0.2</b>	<b>0.666</b>	<b>138.1</b>
32	<b>0.2</b>	<b>0.711</b>	<b>137.9</b>
Avg:	<b>12.65</b>	<b>0.5175</b>	<b>155.67</b>

check CERES estimates with Hc= 12±3km



# Summary

1. Based on observations, tropical regions may be separated into moist and dry regimes although quantitative cutoff between the two is still an open question. It seems more difficult to identify cloudy moist and clear moist.
2. During January to August 1998, the tropical anvil clouds over oceans decrease with increase in SST ( $\sim 25\%/\text{K}$ ).
3. The CERES observed LW and SW fluxes generally differ from Lindzen et al.  $10 \sim 25 \text{ W/m}^2$ , except SW in cloudy moist area where the value is much larger ( $\sim 64 \text{ W/m}^2$ ). The striking feature is that all CERES observed net changes from those of Lindzen et al. point to the same positive feedback direction for the anvil decrease.



## Summary (cont.)

4. Observations and consistency estimations suggest that Lindzen et al. may use a too small albedo for anvil clouds.
5. Using CERES observed radiative properties both 3.5-box greenhouse model calculation and perturbation analysis show that the decrease of anvil clouds with increasing SST results a slightly positive feedback compared to the strong negative feedback of Lindzen et al.